

## **HICO On-Line Atmospheric Correction**



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HyspIRI Aquatic Group June 5, 2015





- HICO the Hyperspectral Imager for the Coastal Ocean Operated on the ISS 2009 2014.
- HICO calibrated radiances
  - On orbit maintenance of calibration
- Atmospheric correction
  - Hyperspectral challenge
  - ATREM, Tafkaa 6S, Tafkaa Tabular
- OSU implementation of Tafkaa 6S
- Examples
  - San Francisco Bay
  - Monterey Bay
  - Lake Erie

Special thanks to NRL, ONR, the DoD Space Test Program, JAXA and NASA for their support of the HICO program.





- HICO fully calibrated in the laboratory (Lucke et al, 2011)
  - Radiometric calibration
  - Spectral calibration
  - Dark current correction
  - Second Order correction
- HICO does not have a second order filter or an on-board calibrator.
- Cannot ask the ISS to rotate to point at the moon.
- On-orbit calibrations using natural scenes (Gao et al, 2012)
  - Spectral calibration using
     Fraunhofer lines and oxygen line
  - Radiometric calibration using land calibration targets
  - Second order correction using water scenes





## **Calibrated Spectral Radiances**





Left: Spectra extracted from pixels along the east-west transect yellow. shown in Approximate locations the of spectra are indicated by same color Xs on the image. Spectra scaled are calibrated at-sensor radiances.

**Right:** Mean and standard deviation of 1295 pixels in the red Region of Interest. SNR The (μ/σ including all sensor and environmental variations) is >300:1 much of the for spectra. Spectra are scaled calibrated atsensor radiances.





# Radiometric Comparison of HICO to MODIS (Aqua)

Nearly coincident HICO and MODIS images of turbid ocean off Shanghai, China demonstrates that HICO is well-calibrated



MODIS (Aqua) Date: 18 January 2010 Time: 05:00:00 UTC Solar zenith angle: 52° Pixel size: 1000 m

#### East China Sea off Shanghai



R.-R. Li and B.-C. Gao, NRL

#### Top-Of-Atmosphere Spectral Radiance



Chlorophyll Comparison of HICO to MODIS (Aqua)



Nearly coincident MODIS and HICO<sup>™</sup> images of the Yangtze River, China taken on January 18, 2010. Left, MODIS image (0500 GMT) of Chlorophyll-a Concentration (mg/m3) standard product from GSFC. The box indicates the location of the HICO image relative to the MODIS image. Right, HICO<sup>™</sup> image (0440 GMT) of Chlorophyll-a Concentration (mg/m3) from HICO<sup>™</sup> data using ATREM atmospheric correction and a standard chlorophyll algorithm. (R-R Li and B-C Gao, NRL)





signal

Atmosphere most of

- Atmospheric gases are well mixed, well understood
- Water is variable but well understood
- Aerosols variable in space & time
- Accurate aerosol models and radiative transfer necessary



Multispectral channels selected to avoid water vapor and other absorptions
 Must correct the full spectrum for hyperspectral data

Figure From Menghua Wang, NOAA/NESDIS/STAR





- Tafkaa-6S (Available for HICO data on the OSU HICO website)
  - Based on ATREM (Gao & Davis 1997 PROC SPIE)
  - Uses 6-S atmospheric model
  - User selects aerosol model and optical depth
  - Handles data from all altitudes
  - Uses absorption features to correct water vapor at all wavelengths
  - Changes from ATREM include ability to parse image header file, improved speed, uses larger set of aerosol models
- Tafkaa-Tabular (not generally available, needs further support)
  - Much of the code based on ATREM (Gao & Davis 1997, PROC SPIE)
  - Uses a large look-up table for the aerosol correction
    - Table created using Zia Ahmed's full vector radiative transfer model
    - Can use dark pixel assumption for open ocean scenes
  - Includes a correction for reflections off of the sea surface
  - Only works for near sea-level data
  - Originally described in (Gao, Montes, Ahmad, & Davis, Applied Optics 2000), modifications in several SPIE proceedings

# HICO atmospheric correction tool

- OSU HICO website tool
- uses Tafkaa 6S

### **Tool access types**

- 1) public access
- from "Image Galleries"
- select scenes only
- 2) full access (*available on* request)
- from "Search Data" results
- all scenes



#### **HICO HICO** atmospheric correction sequence



Tiburon_CA	13	273 201	3-04-18 23:24:23 GMT			
output parameter	rrs (remote-sensing reflectance	e) 🔻 📀				
MAJOR	The parameters in the section below have a significant effect on the magnitude of the output.					
aerosol model	maritime 🔻 😮					
aerosol optical depth (tau_550)	0.05					
elevation	0 km 😮					
offset removal	remove positive offset over water 😧					
MINOR	The parameters in the section below make only minor changes to the spectrum; the default values are recommended.					
atmospheric model	automatic 🗸 🖉					
ozone	-1 ppm 🕑					
atmospheric gases	✓ H <sub>2</sub> O ✓ O <sub>3</sub> ✓ NO <sub>2</sub> ✓ O <sub>2</sub>					
water vapor lines						
9	adjacent window	H <sub>2</sub> O band center	adjacent window			
Set	1: 0.705 µm 3 ▼	0.725 µm 5 ▼	0.745 µm 3 ▼			
Set	2: 0.805 µm 3 ▼	0.825 µm 5 ▼	0.845 µm 3 ▼			
process						
to L2						

1. On the OSU HICO website, find a HICO scene and click

> process to L2

**2.** Select appropriate atmospheric values for the scene.

3. Examine the spectra at userselected locations within the scene. Reprocess as needed, then download the data.









#### **HICO Atmospheric Correction**

The atmospheric correction algorithm used is tafkaa\_6s from NRL, developed by Marcos Montes, Bo-Cai Gao, and Curtiss Davis. Further details are available in the <u>Tafkaa User's Guide</u>.

Monterey_Bay	y_CA	6411	2011-05	5-03 21:50:12 GMT			
output parameter	rrs (	remote-sensing reflecta	nce) 🗸	0			
MAJOR	The p outpu	The parameters in the section below have a significant effect on the magnitude of the output.					
aerosol model	mari	itime 🧹 😮					
aerosol optical depth (tau_550)	0.2	0.2 0					
elevation	0	0 km 🙆					
offset removal	🗌 re	remove positive offset over water 😧					
MINOR	The p defau	The parameters in the section below make only minor changes to the spectrum; the default values are recommended.					
atmospheric model	automatic 🗸 🗸						
ozone	-1	-1 ppm 2					
atmospheric gases		☑ H <sub>2</sub> 0 ☑ 0 <sub>3</sub> ☑ NO <sub>2</sub> ☑ 0 <sub>2</sub> ⊘					
water vapor lines							
	8	adjacent window	H <sub>2</sub> O band center	adjacent window			
	Set 1:	0.705 µm 3 🗸	0.725 µm 5 💙	0.745 µm 3 🗸			
	Set 2:	0.805 µm 3 🗸	0.825 µm 5 🗸	0.845 µm 3 🗸			
process to L2							

The tool includes two sections that have Major and Minor Effects. Major: Aerosol Model Aerosol Optical Depth Elevation Offset Removal (assume Rrs from 740-785 nm = 0)

#### Minor:

#### Atmospheric Model (automatically selected for location and season) Ozone (use climatology or input a value from your data) Water Vapor (uses the selected absorption lines)



## **Atmospheric Correction Help Buttons**



output parameter	rrs (remote-s	ensing reflectance)	<b>~ ⊘</b>		
MAJOR	The parameter output.	s in the section below	have a significant effect or	n the magnitude of the	
aerosol model	maritime N	<ul> <li>2</li> </ul>			
aerosol optical depth (tau_550)	0.2 🚱				
	The aerosol optical depth (AOD) has a large effect on the resulting atmospheric correction. If the AOD value is too low, it can cause the spectra in the red to be above zero (unrealistic). If the AOD value is too high, it can cause negative radiances and reflectances. The default is 0.2. The range is approximately 0 - 2.0.				
elevation	0	km 😮			
offset removal	remove po	sitive offset over water	, <b>0</b>		
MINOR	The parameters in the section below make only minor changes to the spectrum; the default values are recommended.				
atmospheric model	automatic V				
ozone	-1 ppm 😮				
atmospheric gases	♥ H <sub>2</sub> 0 ♥ 0 <sub>3</sub> ♥ NO <sub>2</sub> ♥ 0 <sub>2</sub> ❷				
water vapor lines					
	e adja	acent window	H <sub>2</sub> O band center	adjacent window	
	Set 1: 0.7	705 µm 3 ❤	0.725 µm 5 💙	0.745 µm 3 💙	
	Set 2: 0.8	05 µm 3 💙	0.825 µm 5 🗸	0.845 µm 3 💙	

There is a help button for each topic; click on it to get help and then click again to remove the help screen. HICO image: Tiburon, California (2014)



HICO

Remote sensing reflectance spectra from February 19, 2014 for: (1) Suisun Bay, (2) San Pablo Bay, and (3) offshore showing outflow from the

bays even during low flow drought conditions.

# HICO image: Tiburon, California (2011)

**HICO** 

### HICO provides 90 m GSD, hyperspectral data, and high SNR



#### HICO Atmospheric Correction for Monterey\_Bay\_CA



# Additional tools to examine your data:

Image showing negatives Move the circles to any location in the image Change scales and reset the spectra Across Scene Transect Includes one selected channel.

Using 410 highlights an issue with enhanced blue at the scene edges for some open ocean scenes



## Microcystis bloom in Lake Erie





HICO Image of a massive *Microcystis* bloom in western Lake Erie, September 3, 2011 as confirmed by spectral analysis.





### Hyperspectral Imager for the Coastal Ocean (HICO)





- Built and launched in 28 months
- Installed September 24, 2009 on the ISS

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- > Operated for 5 years
- > 10,000 scenes
- > 50 projects

**HICO** 

OSU Tafkaa 6S atmospheric correction

Data from NASA: <u>http://oceancolor.gsfc.nasa.gov</u> and at: <u>http://hico.coas.oregonstate.edu</u>





## Backup





The apparent reflectance  $\rho^*_{obs}$  at a hyperspectral sensor for a given wavelength is

$$\rho_{obs}^* = \pi L_{obs} / (\mu_o F_o) \tag{1}$$

where  $L_{obs}$  is the radiance of the ocean–atmosphere system measured by the sensor,  $\mu_o$  is the cosine of the solar zenith angle, and  $F_o$  is the extraterrestrial downward solar irradiance at the top of the atmosphere. Then  $\rho_{obs}^*$  can be expressed as:

$$\rho_{obs}^* = T_g \left[ \rho_{atm+sfc}^* + \rho_w t_d t_u / (1 - s\rho_w) \right]$$
(2)

where  $T_g$  is the total atmospheric gaseous transmittance on the sun-surface–sensor path,  $\rho^*_{atm+sfc}$  is the reflectance resulting from scattering by the atmosphere and specular reflection by ocean surface facets,  $t_d$  is the downward transmittance (direct + diffuse),  $t_u$  is the upward transmittance, *s* is the spherical albedo that takes into account the reflectance of the atmosphere for isotropic radiance incident at its base, and  $\rho_w$  is the water- leaving reflectance. Solving (2) for  $\rho_w$  yields

$$\rho_w = \rho_{obs}^* / T_g - \rho_{atm+sfc}^* / [t_d t_u + s (\rho_{obs}^* / T_g - \rho_{atm+sfc}^*)]$$
(3)

Given  $L_{obs}$ , the water-leaving reflectance can be derived according to (1) and (3) and the other quantities in the right hand side of (3) modeled theoretically.