



Imaging Spectroscopy of the Global Cryosphere

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Science Questions

What are the controls on the global snowalbedo feedback?

What is the response of global snow and ice albedo and melt to temperature increases and dust/BC radiative forcing?



WORKING GROUP I CONTRIBUTION TO THE FIFTH ASSESSMENT REPORT OF THE INTERGOVERNMENTAL PANEL ON CLIMATE CHANGE









Aren't we already dealing with these with ICESat, ICEBridge, and ICESat-2?!

Nope – those are just the changes in mass – no attribution to forcing of changes in mass.



Background



- The snow albedo feedback amplifies changes in global radiative forcing, particularly in high latitude regions which are highly sensitive to changes in climate
- Snow/ice albedo and melt are highly sensitive to trace amounts of light absorbing impurities (e.g. black carbon)
- Roughly one-fifth of Earth's population uses melted snow or glacial ice in some capacity for water resources
- Glacier and ice sheet runoff accounts for a large fraction of sea level rise
- The timing and magnitude of snowmelt runoff modulate regional climate and hydroclimate, dominates ecosystem function in vast regions, and contribute to most major flooding events

Uncertainties in our understanding of the processes that control snow albedo feedbacks result in large uncertainties in projected changes in climate and sea level



Global Cryosphere





$\frac{dU}{dt} + Q_m = (1 - \alpha)S + L^* + Q_s + Q_v + Q_g + Q_r$



Elk Range, Colorado River Basin, April 2009



Greenland Melt

JPL

The Cryosphere, 6, 821–839, 2012 www.the-cryosphere.net/6/821/2012/ doi:10.5194/tc-6-821-2012 © Author(s) 2012. CC Attribution 3.0 License.





Greenland ice sheet albedo feedback: thermodynamics and atmospheric drivers



Fig. 1. Melt extent over the GIS determined from Oceansat-2 satellite scatterometer, Special Sensor Microwave Imager/Sounder, and Moderate-resolution Imaging Spectroradiometer satellite data for (*A*) July 8, 2012, and (*B*) July 12, 2012. Red areas indicate melt detected by the satellites, white areas indicate no melt, and blue represents ocean. The surface of almost the entire ice sheet, including the dry snow region, experienced melt on July 12, 2012. Figure courtesy of Dorothy Hall, NASA Goddard Space Flight Center.

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Greenland Albedo

Climate change and forest fires synergistically drive widespread melt events of the Greenland Ice Sheet

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0.01 30E 30E 30W 30W

Figure 2. Spatial distribution of CAM5 (IMPRV) simulated ten-year monthly mean (April) (a) BC concentration in top snow layer (in ng g^{-1}) and (b) forcing induced by BC in snow column (W m⁻²), averaged when snow is present over land.





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Hindu Kush-Himalaya



38°0'0"N

36°0'0"N

34°0'0"N

2°0'0"N



BC ends LIA in Alps

End of the Little Ice Age in the Alps forced by industrial black carbon

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Glaciers in the European Alps began to retreat abruptly from their mid-19th century maximum, marking what appeared to be the end of the Little Ice Age. Alpine temperature and precipitation records suggest that glaciers should instead have continued to grow until circa 1910. Radiative forcing by increasing deposition of industrial black carbon to snow may represent the driver of the abrupt glacier retreats in the Alps that began in the mid-19th century. Ice cores indicate that black carbon concentrations increased abruptly in the mid-19th century and largely continued to increase into the 20th century, consistent with known increases in black carbon emissions from the industrialization of Western Europe. Inferred annual surface radiative forcings increased stepwise to 13-17 $W \cdot m^{-2}$ between 1850 and 1880, and to 9–22 $W \cdot m^{-2}$ in the early 1900s, with snowmelt season (April/May/June) forcings reaching greater than 35 W·m⁻² by the early 1900s. These snowmelt season radiative forcings would have resulted in additional annual snow melting of as much as 0.9 m water equivalent across the melt season. Simulations of glacier mass balances with radiative forcingequivalent changes in atmospheric temperatures result in conservative estimates of accumulating negative mass balances of magnitude -15 m water equivalent by 1900 and -30 m water equivalent by 1930, magnitudes and timing consistent with the observed retreat. These results suggest a possible physical explanation for the abrupt retreat of glaciers in the Alps in the mid-19th century that is consistent with existing temperature and precipitation records and reconstructions.

Simulations of glacier-length variations using glacier flow and mass balance models forced with instrumental and proxy temperature and precipitation fail to match the timing and magnitude of the observed late 19th century retreat (8–12). Matches between simulations and observations have only been achieved when additional glacier mass loss is imposed after 1865, or when

preci Significance

treat

13). The end of the Little Ice Age in the European Alps has long riables been a paradox to glaciology and climatology. Glaciers in the obsel Alps began to retreat abruptly in the mid-19th century, but

Hι reconstructions of temperature and precipitation indicate that veglaciers should have instead advanced into the 20th century. rd tweet We observe that industrial black carbon in snow began to inof the crease markedly in the mid-19th century and show with simsumn ht. ulations that the associated increases in absorbed sunlight by that. is black carbon in snow and snowmelt were of sufficient magnot f nitude to cause this scale of glacier retreat. This hypothesis ulatior offers a physically based explanation for the glacier retreat that Ve mass maintains consistency with the temperature and precipitation to reconstructions. point ıre

not captured well in the ambient climatic records" [italics added for emphasis] (10). In other words, Huybrechts et al. (10) suggest that some forcing beyond changes in temperature and precipitation was driving a large part of the negative mass balance.

Glacier mass balance is controlled by accumulation at higher elevations and ablation at lower elevations. In each year, glacier

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2,



Dust RF in Colorado River Basin

WATER RESOURCES RESEARCH, VOL. 48, W07522, doi:10.1029/2012WR011986, 2012

Dust radiative forcing in snow of the Upper Colorado River Basin: 2. Interannual variability in radiative forcing and snowmelt rates

S. McKenzie Skiles,^{1,2} Thomas H. Painter,^{1,2,3} Jeffrey S. Deems,^{4,5} Ann C. Bryant,⁶ and Christopher C. Landry⁷



Response of Colorado River runoff

to dust radiative forcing in snow





How fast rivers come up





Uncertainties in Glacier Mass Balance contributions to SLR

Antarctic & Subantarctic	Α	B H
High Mountain Asia		
Arctic Canada North		
Alaska		
Greenland		
Russian Arctic	التنبي ا	
Arctic Canada South		
Svalbard	Here Here	
Southern Andes	Freeday A	
Western Canada /US		Her H
Iceland	HAR I	124
Low Latitudes		
Scandinavia	Hel •	
North Asia		*
Central Europe		
New Zealand		• GRACE
Caucasus & Middle East		ICESat
	-2000 -1000 0 1000 2000	-100 -80 -60 -40 -20 0
	mass budget [kg m ⁻² yr ⁻¹]	mass budget [Gt yr ⁻¹]

Gardner et al 2013

Why a spectrometer?





Robust retrievals that come from imaging spectrometer and provide critical understanding of cryosphere forcings

- ✤ Spectral albedo
- Snow grain size
- Dust/BC radiative forcing in snow and ice
- Snow surface liquid water content

All retrievals are demonstrated with AVIRIS or ASO



Spectral resolution drivers

- To understand changes in snow depth/ice height to within 0.25 m
- With enthalpy of fusion, a 0.25 m snowpack with density 80 kg/m³ is equivalent to 6 W/m² or instantaneous 15 W/m²
- \rightarrow < 40 nm spectral resolution





Again, the Science Questions

What are the controls on the global snow-albedo feedback? What is the response of global snow and ice albedo and melt to temperature increases and dust/BC radiative forcing?

We need a spaceborne imaging spectrometer to answer these critical questions.

The rest of the water cycle has or will soon have missions.

Such a mission directly addresses the fundamental questions in Climate Change and Variability, Water and Energy Cycle, Carbon Cycles and Ecosystems through understanding the cryosphere



Airborne Snow Observatory

Setting the stage



Albedo

A

Imaging Spectrometer 0.35-1.05 μ m 2 m spatial resolution from 4000 AGL

RIEG

30 Scanning LiDAR 1064 nm 1 m spatial resolution Snow Water Equivalent

۸S



ASO Spectral Albedo







ASO-DM1 Results







ASO-DM1 Results







ASO-DM1 Results









ASO SWE and albedo 2014



Snow Water Equivalent 2014



Snow Albedo 2014



ASO Tuolumne snowmelt 2013





Why is it urgent

- Relative contributions by forcings of mass loss from glaciers and ice sheets remain unknown due insufficient measurements and constraints on regional climate and hydroclimate modeling
- Mountain glaciers and ice sheets have ongoing retreat and downwasting
- A strong mitigation opportunity against warming impacts on the cryosphere may rest in reducing particulate loading – unlike reducing CO₂ a "regional solution to a regional problem"
- The World Bank is calling for these mitigation efforts already, but measurements are needed now to be able to understand efficacy of mitigation effort.



A JOINT REPORT OF The World Bank The International Cryosphere Climate Initiative

