

Deriving soil moisture and sediment mobility using future HyspIRI-derived thermal inertia





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Scheidt, S., Ramsey, M. & Lancaster, N., Determining soil moisture and sediment availability at White Sands Dune Field, NM from apparent thermal inertia data, *JGR – Earth Surface, in press,* 2009.



HyspIRI Relevant Science

- TQ3: Water Use and Availability
 - Can we improve early detection, mitigation, and impact assessment of droughts at local to regional scales anywhere on the globe?
 - How does the partitioning of precipitation into ET, surface runoff and ground-water recharge change during drought?

CQ5: Surface Composition and Change

 How is the composition of exposed terrestrial surface responding to anthropogenic and non anthropogenic drivers (e.g., desertification, weathering, climate change, human disturbance)?

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What Is Missing in the CQ's?

- Thermophysical Properties
 - interpret diurnal, seasonal, or climatic surface and subsurface temperature variations
 - <u>thermal inertia</u>: controls the amplitude of the daily thermal curve, I = (kρc)^{1/2}
 - <u>albedo:</u> controls the average temperature fluctuation
 - <u>ATI:</u> apparent thermal inertial, $ATI = (1-A)/(T_d-T_n)$
 - derived products:
 - soil moisture (linkage to future SMAP data)
 - particle size, density
 - soil erosion potential in arid/semi-arid lands
 - > direct relationship to surface roughness/soil moisture

Background: Soil Moisture

- Measurements difficult in large arid lands
 - permeable sands typically experience a quick drying
 - field measurements over a large geographical region are difficult and expensive
 - remote location and difficult working conditions
 - the need for very sensitive in-situ sensors to measure the low water / high salinity content of desert soils
 - soil mobility
 - important for desertification/agriculture conditions
 - Iocal to regional climate impacts
 - the upper few centimeters are the most critical for understanding erosion potential

Background: Soil Moisture

- Using Microwave Sensors
 - regional to continental scale (e.g., PSR, AMSR-E)
 - future Soil Moisture Active Passive (SMAP)
 - spatial resolution of 1-3 km every 12 hours
 - surface roughness and vegetation
 - > important for sediment transport studies
 - spatial resolution of 25 km (AMSR-E) and 1-3 km (SMAP) limits the interpretation of soil moisture for specific landforms
 - example, the White Sands eolian system is only 26 km²
 - ~1 AMSR-E pixel, ~80 SMAP pixels
 - ~190,000 HyspIRI pixels

Background: White Sands, NM

- Understand the relationship between
 - apparent thermal inertia (ATI) ↔ thermal inertia ↔
 surface soil moisture ↔ sediment availability

• Objectives:

- how are the differences in ATI over time explained for the same geographic areas?
 - can these be used to determine soil moisture/sediment mobility?

 how do the calculations of ATI compare between ASTER and MODIS Terra?



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Background: White Sands, NM

Eolian System

- gypsum dominated
- material is transported by W winds from ephemeral playa lakes
 - timing is determined by wet-dry cycles
 - resupply determined by the availability of the gypsum created by evaporation and mineral formation



area of overlap of the ASTER day and night images (green arrow - Lake Lucero)



Background: Thermal Inertia

Previous studies

- have used thermal inertia to estimate soil moisture
 - Price (1985), Xue (1986), Zhang et al. (2002), Zhenhua & Yingshi (2006), Cai et al. (2007)

New work

[Scheidt et al., 2009]

- reproduced a thermal inertia model for ASTER data Xue and Cracknell (1995)
 - programmed in IDL/ENVI
 - White Sands data are now presented in actual thermal inertia units (TI) units rather than ATI
 - TI is related to soil moisture and sediment availability



Models



Soil moisture vs. thermal inertia as a function of soil density reproduced from lookup (Mae and Xue, 1990)



Fécan et al. (1999) show the erosion threshold velocity ratio as a function of soil moisture

the Mae and Xue (1990) look up tables have been combined with the Fécan et al. (1999) relationship



$$P = \frac{(1-a)S_0C_t}{\Delta T\sqrt{\omega}} \left\{ \frac{A_1 \left[\cos(\omega t_2 - \delta_1) - \cos(\omega t_1 - \delta_1) \right]}{\sqrt{1 + \frac{1}{b} + \frac{1}{2b^2}}} + \frac{A_2 \left[\cos(\omega t_2 - \delta_2) - \cos(\omega t_1 - \delta_2) \right]}{\sqrt{2 + \frac{\sqrt{2}}{b} + \frac{1}{2b^2}}} \right\}$$

Data Inputs

a = albedo

 ΔT = temperature difference

 $t_n = day (n=1) and night (n=2) time$

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 $\mathbf{S}_0 = \text{solar constant}$

- $\omega = angular \ velocity \ of \ Earth$
- $C_t = atmospheric transmissivity$
- $\delta_n = phase difference^*$
- A_n = coefficient of Fourier series*
- * values are dependent on solar declination, latitude, and maximum daytime temperature

Data Output

Xue and Cracknell (1995)

P: thermal inertia (J m⁻² K⁻¹ s^{-1/2} or TIU) *ATI*: (1-a)/⊿T

Sources of Error

Albedo: dependent on instrument and calculation

▲ T: satellite overpasses are not at maximum and minimum diurnal temperatures

C_t: assumed to be constant or corrected

time: overpasses are 36 hours apart (potential error and in-scene variability, i.e., cloud, rain, etc.)



- Examined ASTER and MODIS apparent thermal inertia (ATI)
 - calculated over 8 years at White Sands, NM
 - 54 (day) and 70 (night) cloud-free ASTER images
 - seven day-night image pairs were selected
 - closest day/night pairs (36 hour difference)





ASTER Data

• ASTER (2000-2008)

- all data were processed to Level 2 (surface reflectance and kinetic temperature)
- focus on two time pairs
 - > April 2006 (dry period) and April 2007 (wet period)

Related Image and Weather Station Statistics				
Image Pair		Image	Prior 30-day	%Area Playa
Date	ΔT air (K)	ΔT (K)	Precip (mm)	Inundation
Feb 7-8, '02	ND	16.8	ND	15
Nov 6-7, '02	27	22.1	26	0
Nov 22-23, '02	27	17.3	21	0
May 3-4, '04	25	33.5	21	0
April 7-8,'06	23.8	27	0	0
April 26-27, '07	22.2	32	15	50
March 11-12,'08	ND	26.5	ND	0



ASTER VNIR Data

27 April 2007 (wet period)

8 April 2006 (dry period)





ASTER: 8 April 2006

albedo







ASTER Relative Albedo Changes





ASTER Derived Thermal Inertia





ASTER Derived Soil Moisture



range of soil moisture values (9% - 25%)



ASTER Derived Soil Moisture





ASTER Derived Soil Moisture









ASTER Erosion Ratio



time-averaged image of the unit less erosion threshold wind velocity ratio (u_{θ}^*/u_{d}^*)



MODIS: 14 March 2008





GOES: 14 March 2008





- Retrieved TI, Soil Moisture, and Wind Threshold Velocity
 - at a high spatial resolution using ASTER
 - for eolian systems
 - soil moisture is the most important parameter where estimating sediment availability
 - values are slightly higher than would be expected for the dunes
 - > model needs further refinement and testing in other regions
 - bowever, the general trends within the dunes and over time are valid
 - > detected a dramatic decrease in soil moisture in March 2008
 - several days before the largest dust storm at White Sands in over a decade (used as a monitoring tool with HyspIRI?)



extra slides



MODIS Albedo

- MODIS products
 - MOD43:16-day averaged albedo product
 - not retrieved for White Sands
 - false snow or cloud detection?
 - MOD09: reflectance product
 - can be used if assumed to be lambertian surface
 - solution: scale the ASTER albedo to MODIS albedo using scene statistics of the same pixels







Albedo Calculation

- Compared ASTER and MODIS
 - albedo derived from standard reflectance products
 - broadband albedo calculations from Liang (2000)
 - published for both MODIS and ASTER
 - ASTER albedo model:
 - $a = 0.484*b_1 + 0.335*b_3 0.324*b_5 + 0.551*b_6 + 0.305*b_8 0.367*b_9 0.0015$

- MODIS albedo model: • $a = 0.160^{*}b_1 + 0.291^{*}b_2 - 0.243^{*}b_3 + 0.116^{*}b_4 + 0.291^{*}b_2 + 0.291^{*}b_3 + 0.291^{*}b_3 + 0.291^{*}b_3 + 0.291^{*}b_4 + 0.291^{*}b_2 + 0.291^{*}b_3 + 0.291^{*}b_3 + 0.291^{*}b_4 + 0.291^{*}b_3 + 0.291$

0.112*b₅ - 0.081*b₇ - 0.0015